

Abundance of Cadmium in Paleozoic Limestones in Japan

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Cadmium is significantly enriched in Paleozoic (Permian and Carboniferous) limestones in Japan. The geometric mean of Cd in these limestone samples ($n=111$) is 0.57 ppm, which is an order of magnitude higher than that of Quaternary ones. Most of divalent heavy metals in these limestones exist in accessory minerals such as silicates, whereas Cd is contained in calcite. Submarine volcanoes which were active during Paleozoic period might be responsible for the high Cd content of Japanese limestones of this age.

The fixation of Cd^{2+} in calcite is probably an important mechanism in regulating the concentration of this ion in seawater, since the ionic radius (1.09 \AA^{11} in 6-fold coordination) of Cd^{2+} is almost identical with that (1.14 \AA^{11}) of Ca^{2+} . In fact, remarkable incorporation of Cd^{2+} into calcite was observed in laboratory experiments.^{2–5} Limestone of marine origin may therefore serve as a potential reservoir for Cd. However, the abundance of Cd in limestone is not well established due

to lack of data, although much information on this metal in other geological materials is available in connection with environmental pollution.

Mason and Moore⁶ proposed 0.09 ppm as the abundance of Cd in carbonate rocks in 1982. Recently, we reported similar Cd contents (geometric mean: 0.071 ppm, $n=94$) for Quaternary limestones in southwest Japan.⁷ These values are quite similar to the abundance of this element in the continental crust (0.098 ppm)⁸ or igneous rocks (0.08 ppm).⁶

On the other hand, markedly high Cd contents (av. ca. 0.5 ppm; max. 2.4 ppm) have been found in Permian and Carboniferous limestones from several localities in Japan.^{9–12} Large scale deposits of Japanese limestones often occur during the Carboniferous to the Permian stage.¹³ We therefore aimed to confirm the difference in Cd content between Japanese Paleozoic and Quaternary limestones by using new data and to discuss the mechanism of incorporation of this element into these ancient limestones.

Experimental

Twelve sampling localities are shown in Fig. 1, together with those of other samples previously analyzed by the authors.^{7,10–12,14} Yoron-jima samples were collected from the

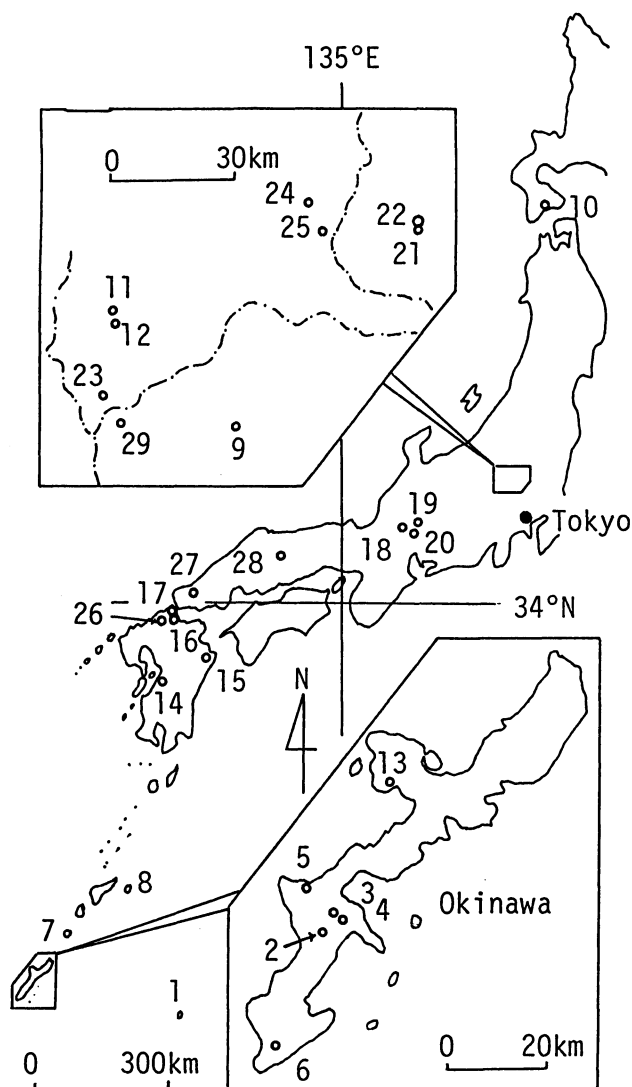


Fig. 1. Sampling localities.

| | | | |
|---------------------|--------|------------------|-------|
| 1. Minamidaito-jima | (Q) | 16. Tsunemi | (P) |
| 2. Itoman | (Q) | 17. Yotsudaka | (P) |
| 3. Akamichi | (Q) | 18. Ibuki | (P) |
| 4. Chibana | (Q) | 19. Takaradani | (P) |
| 5. Okinawa City | (Q) | 20. Akasaka | (P) |
| 6. Yomitan | (Q) | 21. Kuzuu(Upper) | (P) |
| 7. Yoron-jima | (Q*P?) | 22. Kuzuu(Lower) | (P) |
| 8. Kikai-shima | (Q) | 23. Ueno | (P) |
| 9. Chichibu | (T) | 24. Kurohone | (P) |
| 10. Garo | (T) | 25. Kiryu | (P) |
| 11. Aokura | (T-P) | 26. Kawara | (P-C) |
| 12. Shimonita | (T-P) | 27. Ofuku | (P-C) |
| 13. Motobu | (P) | 28. Niimi | (P-C) |
| 14. Ashikita | (P) | 29. Oku-Chichibu | (P-C) |
| 15. Tsukumi | (P) | | |

Italic: Samples analyzed in this work.

Q: Quaternary, Q*P(?): Quaternary and Permian(?), T: Triassic, T-P: Triassic or Permian, P: Permian, P-C: Permian or Carboniferous.

southern shore of this island, and Takaradani samples were from a road cut. The samples of other 10 localities were taken from pits of working or abandoned mines. The geological ages of these samples were identified as the Triassic, Permian, and Permian or Carboniferous.¹⁵⁻¹⁷⁾ The Ritcho Formation from which Yoron-jima samples were taken is non-fossiliferous and estimated to be the Permian.¹⁸⁾ The present Paleozoic limestone deposits are closely associated with basic volcanic rocks.^{17,19)} In addition, they are massive and not stratified. Neither sandstones nor shales are interbedded in these limestone deposits.¹⁹⁾ From these geological features, Horii^{20,21)} proposed a mechanism that submarine volcanism may be responsible for the formation of these limestones.

Samples were pulverized to <0.15 mm in an agate mortar. After the decomposition of samples (2.00 g) with a mixture of nitric and perchloric acids, Cd was determined together with Ni, Cu, and Zn (<5 ppm) by flame atomic absorption spectrometry coupled with a solvent extraction technique using APDC (ammonium pyrrolidinedithiocarbamate). The detection limit of this method is 0.005 ppmCd, and three determinations of a limestone sample give 0.14 ± 0.01 ppmCd. Details of the sample preparation and the analytical procedure were reported previously.¹⁴⁾

Results

The arithmetic means of Cd contents were summarized in Table 1, together with those of some major and trace elements. The prominent feature of the present samples is low SiO₂, Al₂O₃, and Fe contents, indicating that amounts of impurities such as silicates are small. Most of the Ashikita, Tsukumi, and Yotsudaka limestone samples contain a small amount of dolomite. However, X-ray diffractometry for the bulk samples of

other localities gave no peak of other mineral than low-Mg calcite. Discussion was carried out by using the present and our previous data (the sampling localities were given in Fig. 1).

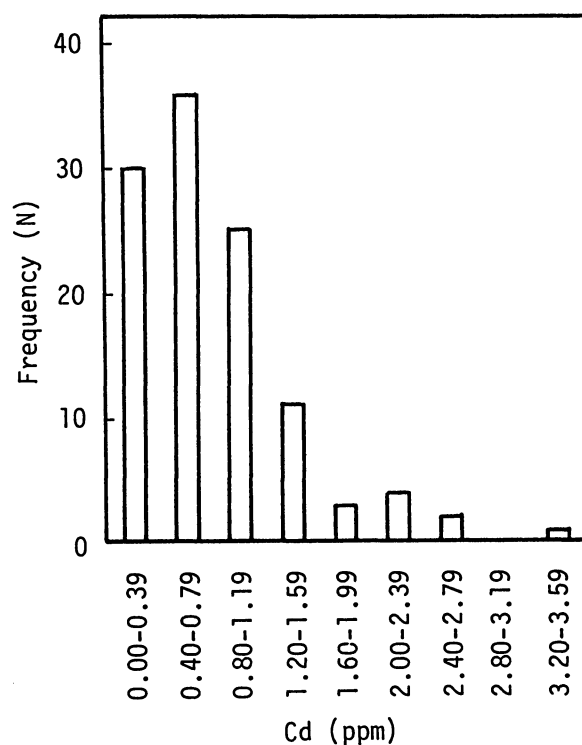


Fig. 2. Frequency distribution of Cd content of Japanese Paleozoic limestone samples.

Table 1. Arithmetic Means of Concentration of Major and Trace Elements in Japanese Limestone Samples (in ppm unless noted as %)

| Sampling locality | No. of samples | SiO ₂ /% | Al ₂ O ₃ /% | CaO/% | MgO/% | Sr | Mn | Fe | Ni | Cu | Zn | Cd |
|-------------------|----------------|---------------------|-----------------------------------|-------|-------|-----|-----|-----|------|-------|------|------|
| Yoron-jima | 8 | 0.06 | 0.014 | 55.28 | 0.34 | 240 | 19 | 160 | 0.42 | 0.27 | 5.1 | 0.42 |
| | | 0.07 | 0.016 | 0.24 | 0.07 | 100 | 12 | 200 | 0.21 | 0.09 | 1.8 | 0.16 |
| Garō | 10 | 0.16 | 0.039 | 54.42 | 0.45 | 500 | 19 | 100 | 0.52 | 0.14 | 2.1 | 0.12 |
| | | 0.13 | 0.001 | 0.79 | 0.24 | 230 | 15 | 59 | 0.22 | 0.08 | 0.9 | 0.04 |
| Ashikita | 9 | 0.37 | 0.020 | 52.03 | 2.01 | 540 | 17 | 93 | 0.70 | 0.22 | 3.2 | 0.49 |
| | | 0.44 | 0.028 | 2.50 | 2.21 | 250 | 11 | 51 | 0.23 | 0.14 | 1.4 | 0.26 |
| Tsukumi | 10 | 0.09 | 0.020 | 53.06 | 1.09 | 210 | 19 | 86 | 0.31 | 0.32 | 4.3 | 0.91 |
| | | 0.07 | 0.015 | 0.71 | 0.52 | 46 | 13 | 55 | 0.20 | 0.06 | 1.8 | 0.42 |
| Tsunemi | 5 | 0.27 | 0.11 | 52.98 | 0.93 | 220 | 27 | 300 | 1.3 | 0.45 | 3.8 | 0.59 |
| | | 0.25 | 0.13 | 0.27 | 0.22 | 29 | 36 | 160 | 1.3 | 0.22 | 1.7 | 0.05 |
| Yotsudaka | 3 | 0.28 | 0.11 | 50.11 | 3.32 | 540 | 350 | 330 | 0.55 | 0.096 | 1.6 | 0.34 |
| | | 0.26 | 0.085 | 1.26 | 0.76 | 400 | 460 | 170 | 0.15 | 0.033 | 0.3 | 0.13 |
| Ibuki | 1 | <0.02 | 0.0025 | 54.51 | 0.85 | 85 | 68 | 130 | 2.0 | 0.93 | 16.1 | 1.1 |
| Takaradani | 4 | 0.10 | 0.0095 | 53.57 | 0.75 | 430 | 19 | 77 | 0.84 | 0.23 | 21.2 | 2.1 |
| | | 0.09 | 0.0065 | 0.42 | 0.13 | 150 | 26 | 91 | 0.45 | 0.17 | 13.2 | 0.77 |
| Akasaka | 10 | <0.02 | 0.0034 | 54.43 | 0.48 | 500 | 18 | 37 | 0.49 | 1.4 | 5.7 | 0.57 |
| | | — | 0.0006 | 0.42 | 0.32 | 370 | 20 | 21 | 0.42 | 0.96 | 4.3 | 0.34 |
| Kawara | 4 | <0.02 | 0.0016 | 54.84 | 0.40 | 170 | 3.6 | 22 | 0.36 | 0.30 | 14.7 | 0.66 |
| | | — | 0.0003 | 0.26 | 0.16 | 40 | 1.9 | 8 | 0.23 | 0.07 | 9.4 | 0.07 |
| Ofuku | 7 | <0.02 | 0.0020 | 54.97 | 0.31 | 320 | 6.7 | 69 | 0.34 | 0.41 | 11.8 | 0.78 |
| | | — | 0.0005 | 0.33 | 0.03 | 400 | 3.3 | 24 | 0.18 | 0.16 | 9.0 | 0.26 |
| Niimi | 5 | 0.21 | 0.0023 | 55.61 | 0.20 | 190 | 8.2 | 160 | 0.17 | 4.8 | 6.0 | 0.48 |
| | | 0.18 | 0.0002 | 0.11 | 0.10 | 90 | 2.0 | 40 | 0.16 | 0.93 | 1.4 | 0.25 |

Italic: Standard deviation.

| | [Paleozoic ($n=111$)] | | | | | | | |
|--------------------------------|-------------------------|--------------------------------|-------|--------|-------|------|------|--------|
| | SiO ₂ | Al ₂ O ₃ | Fe | Mn | Ni | Cu | Zn | Cd |
| SiO ₂ | — | 0.57 | 0.73 | 0.68 | 0.72 | 0.53 | 0.51 | 0.10 |
| Al ₂ O ₃ | 0.79 | — | 0.91 | 0.068 | 0.75 | 0.27 | 0.70 | -0.016 |
| Fe | 0.83 | 0.93 | — | 0.29 | 0.88 | 0.38 | 0.76 | 0.068 |
| Mn | 0.58 | 0.77 | 0.78 | — | 0.39 | 0.44 | 0.19 | 0.27 |
| Ni | 0.82 | 0.89 | 0.96 | 0.73 | — | 0.48 | 0.68 | 0.20 |
| Cu | 0.68 | 0.89 | 0.88 | 0.72 | 0.84 | — | 0.29 | 0.21 |
| Zn | 0.76 | 0.96 | 0.95 | 0.78 | 0.92 | 0.90 | — | 0.39 |
| Cd | -0.10 | 0.042 | 0.018 | -0.037 | 0.058 | 0.18 | 0.10 | — |
| | [Quaternary ($n=94$)] | | | | | | | |

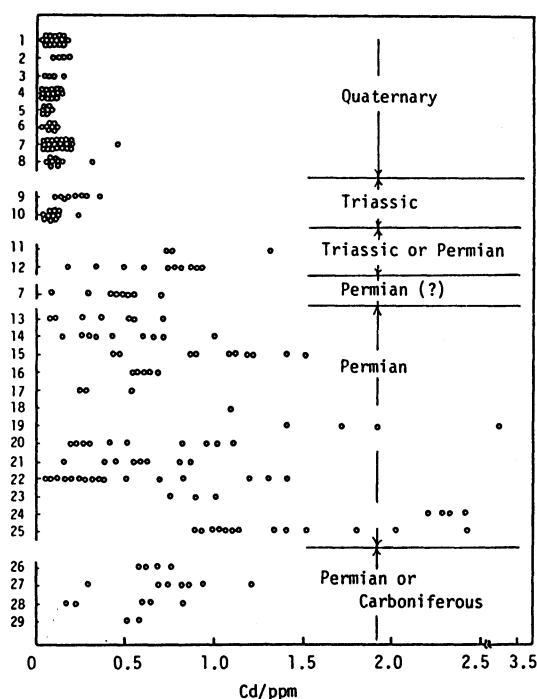


Fig. 3. Cadmium contents of limestone samples from Japan. Numbers on the ordinate correspond to the numbers in Fig. 1.

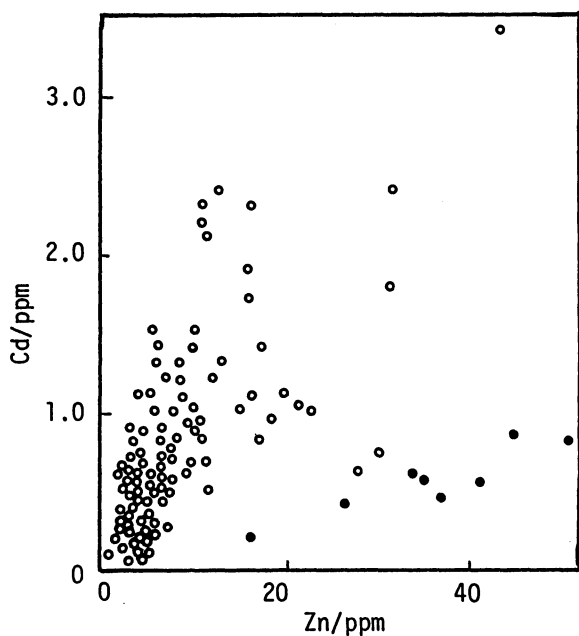


Fig. 4. Relationship between Cd and Zn Contents of Japanese Paleozoic limestone samples.
(●) Kuzuu Lower limestone samples
Except for Kuzuu Lower limestone samples in which an appreciable amount of pyrite is present, correlation coefficient is 0.64 ($n=103$).

It should be noted that the mode of occurrence of Japanese Paleozoic limestones is quite different from those of other regions such as North America, Europe, and China where stratified limestones are prevailing.

Most of Japanese Paleozoic limestones are massive and associated with basic volcanic rocks. This fact suggests that these basic rocks were originated from submarine volcanoes which were highly active during Paleozoic period.¹⁹⁾ Thus, the significant contribution of volcanic activities to high Cd content of these limestones is easily acceptable, because Cd is one of the most enriched elements in volcanic gases or aerosols due to its high volatility²⁶⁻³⁴⁾ (0.01—ca. 1 ppm in gas condensates).

A positive relationship between Zn and Cd contents of Paleozoic limestones (Fig. 4) also suggests the volcanic origin of Cd. Zinc in these limestones is in part present in silicates, at the same time it also occurs in calcite by replacing Ca^{2+} .^{12,35-37)} Zinc contents of volcanic gases and aerosols are one or two orders of magnitude higher than those of Cd,²⁶⁻³⁴⁾ and Cd usually behaves geochemically together with Zn. Therefore the high Zn content of Paleozoic limestones relative to that of Quaternary ones (Table 2) and the above positive relationship could be the evidence for the volcanic origin of Cd. On the other hand, the low Ni and Cu contents of Paleozoic limestones comparable to those of Quaternary ones (Table 2) may be due to the difficulty in substituting Ca^{2+} by these ions in calcite,^{36,37)} although an appreciable amount of Cu^{2+} is contained in volcanic gases and aerosols.^{26,27,29,33,34)}

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References

- 1) R. D. Shannon, *Acta Crystallogr., Sect. A*, **32**, 751 (1976).
- 2) O. Fujino, T. Kumagai, T. Shigematsu, and M. Matsui, *Bull. Inst. Chem. Res., Kyoto Univ.*, Vol. 54, No. 5, 312 (1976).
- 3) Y. Kitano, N. Kanamori, and R. Fujiyoshi, *Geochem. J.*, **12**, 137 (1978).
- 4) R. B. Lorens, *Geochim. Cosmochim. Acta*, **45**, 553 (1981).
- 5) J. A. Davis, C. C. Fuller, and A. D. Cook, *Geochim. Cosmochim. Acta*, **51**, 1477 (1987).
- 6) B. Mason and C. B. Moore, "Principles of Geochemistry," 4th ed, John Wiley & Sons, Inc., New York (1982), p. 176.
- 7) S. Aizawa and H. Akaiwa, *Chikyukagaku (Geochemistry)*, in press.
- 8) H. Heinrichs, B. Schulz-Dobrick, and K. H. Wedepohl, *Geochim. Cosmochim. Acta*, **44**, 1519 (1980).
- 9) T. Sakae, K. Yoshimura, and T. Tarutani, *Bull. Akiyoshi-dai Mus. Natl. History*, No. 16, 1 (1981).
- 10) S. Aizawa and H. Akaiwa, *Chikyukagaku (Geochemistry)*, **21**, 21 (1987).
- 11) S. Aizawa and H. Akaiwa, *Chikyukagaku (Geochemistry)*, **21**, 31 (1987).
- 12) S. Aizawa and H. Akaiwa, *Chem. Geol.*, **98**, 103 (1992).
- 13) T. Hashimoto, "Limestones in Japan," ed by S. Kawada, Sekkaiseki Kogyo Kyokai, Tokyo (1983), p. 95.
- 14) S. Aizawa and H. Akaiwa, *Chem. Geol.*, **67**, 275 (1988).

- 15) S. Sakagami, S. Minamikawa, and M. Kawashima, *Chigaku Zasshi*, **78**, 37 (1969).
 - 16) T. Fujinuki, "Limestones in Japan," ed by S. Kawada, Sekkaiseki Kogyo Kyokai, Tokyo (1983), p. 43.
 - 17) "Limestones in Japan," ed by S. Kawada, Sekkaiseki Kogyo Kyokai, Tokyo (1983), Part 4, pp. 257—485.
 - 18) H. Nakagawa, Tohoku Univ., Inst. Geol. Pal., Contr. No. 63, 1 (1967).
 - 19) K. Kanmera, "Sedimentary Rocks in Japan," ed by S. Mizutani, Y. Saito, and K. Kanmera, Iwanami Shoten, Tokyo (1987), Chap. 3, pp. 123—133.
 - 20) M. Horii, *Sekkaiseki*, **78**, 675 (1962).
 - 21) M. Horii, "Limestones in Japan," ed by S. Kawada, Sekkaiseki Kogyo Kyokai, Tokyo (1983), p. 25.
 - 22) G. Marowsky and K. H. Wedepohl, *Geochim. Cosmochim. Acta*, **35**, 1255 (1971).
 - 23) H. Gong, A. W. Rose, and N. H. Suhr, *Geochim. Cosmochim. Acta*, **41**, 1687 (1977).
 - 24) H. Heinrichs, *Fresenius Z. Anal. Chem.*, **294**, 345 (1979).
 - 25) Y. Kajiware, *Geochem. J.*, **10**, 43 (1976).
 - 26) E. A. Lepel, K. M. Stefansson, and W. H. Zoller, *J. Geophys. Res.*, **83**, 6213 (1978).
 - 27) P. Baut-Menard and M. Arnold, *Geophys. Res. Lett.*, **5**, 245 (1978).
 - 28) J. M. Phelan, D. L. Finnegan, D. S. Ballantine, and W. H. Zoller, *Geophys. Res. Lett.*, **9**, 1093 (1982).
 - 29) I. A. Menyailov, L. P. Nikitina, V. N. Shapar, and A. Z. Miklishansky, *J. Geophys. Res.*, **87**, 11113 (1982).
 - 30) J. B. Gemmell, *J. Volcanol. Geotherm. Res.*, **33**, 161 (1987).
 - 31) B. M. Crowe, D. L. Finnegan, W. H. Zoller, and W. V. Boynton, *J. Geophys. Res.*, **92**, 13708 (1987).
 - 32) R. B. Symonds, W. I. Rose, M. H. Reed, F. E. Lichte, and D. L. Finnegan, *Geochim. Cosmochim. Acta*, **51**, 2083 (1987).
 - 33) M. Pennishi, M. F. Le Cloarec, G. Lambert, and J. C. Le Roulley, *Earth Planet. Sci. Lett.*, **88**, 284 (1988).
 - 34) J. P. Quisefit, J. P. Toutain, G. Bergametti, M. Javoy, B. Cheynet, and A. Person, *Geochim. Cosmochim. Acta*, **53**, 2591 (1989).
 - 35) A. Tsusue and H. D. Holland, *Geochim. Cosmochim. Acta*, **30**, 439 (1966).
 - 36) Y. Kitano, A. Tokuyama, and N. Kanamori, *J. Earth Sci., Nagoya Univ.*, **16**, 1 (1968).
 - 37) H. Akaiwa, H. Kawamoto, S. Aizawa, and M. Kaneko, Annual Meeting of the Geochemical Society of Japan, Hakodate, October 1978, Abstr., No. 5C10.
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